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Analysis of Historical Drought Using Standardized Precipitation Index and Reconnaissance Drought Index (RDI) in Tharparkar, Sindh

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Abstract Drought is recognized as disastrous hazard that can be take place in any terrain and can cause significant harming effects to environment of earth. Thus, the drought assessment is of crucial importance for planning mitigation and water resources management. In this study, historical drought was analyzed in the Tharparkar district of Sindh region of Pakistan. The two drought approaches namely: Standardized Precipitation Index and

Reconnaissance Drought Index (RDI) were employed to assess the drought. The long term 35 years monthly precipitation (mm) and temperature (°C) data were used to calculate SPI and RDI. From the analysis, it was observed that extreme drought noticed by SPI and RDI in the years 1986-1987 and 1999-2000 with the values - 2.48 and -4.63. however, both indicators perceived severe and moderate drought during 20001-20004. The strong correlation (r=0.158) comprehend by RDI. From the findings it is advised that to use RDI and SPI to monitor the drought in the region, and for water resources and mitigation planning.

Keyword Drought monitoring, Standardized Precipitation Index, Reconnaissance Drought Index (RDI)

Introduction

Drought is categorized the dangerous meteorological threat of nature (Mishra & Singh, 2010). It is momentary situation which continues for months to years. Corresponding to Whilhit, drought can appear in any kind of environment and in any region in the world (Wilhite, 2000). Now a day's major cause of drought in the world is climate variability and is also responsible for large economic losses over almost every region every year. Drought also could be defined as a continuous period of insufficient precipitation resulting in extreme damage to crops, and loss of yield (Folger & Cody, 2014). Generally, there are four types of drought, meteorological, agricultural, hydrological, and socio-economical. Meteorological drought occurred from the deficiency of precipitation over regions with respect to their climatology. Agricultural drought assigns to a decline in soil moisture and crop water demand (Mishra & Singh, 2010). Hydrological drought, due to the downturn in the flow of water in rivers, reservoirs, or canals (Dracup, Lee, & Paulson Jr, 1980). Socioeconomic drought appeared by the influence of already stated 3 types of drought conditions at the delivery and requirement of economic goods and offerings such as fish, food grain, forage, and hydroelectric power (Marshall & Zhou, 2004). Recently, a fifth class of drought has been notified as to the drought of groundwater (Mishra & Singh, 2010).

Meteorological drought accumulates the agriculture drought as it materializes due to deficiency of rainfall and soil moisture throughout the growing season to sustenance vigorous growth of crop up to adulthood, causing crop stress and abating. It is considered when an area receives less than 75% seasonal rainfall of its normal lasting value (Indian Meteorological Department - IMD). Further, it is categorized as temperate and severe drought if the precipitation deficit is 26–50% and rainfall shortfall exceeds 50% of normal, respectively. A year is taken into account to be a dearth year for the country if the drought exaggerated region exceeds 20% of the entire area of the country (Karinki & Sahoo, 2019). In recent years, floods and drought have befallen with better depth and harshness stages. Drought presently is accountable for increasing mortality rate and economic losses in the world. The drought has a vast effect on bad economies than on extra evolved ones. for example, in many nations of western Africa, 30% population is unprotected to drought each year the income of peoples robustly threatened (Rothman, van Bers, Bakkes, & Pahl-Wostl, 2009). In China and India, averagely 0.5 million people have affected annually through drought economically and environmentally. In 2011, 12 million people were affected in the entire east African region, and food crises occurred across Djibouti, Ethiopia, Kenya, and Somalia. Europe is also distressed from 1998-2009 through a drought in the result of forest fires and heatwaves 80,000 peoples were stricken and overall losses € 4940 billion were estimated(Agency, 2010). Australia has also experienced agricultural and industrial drought from 1995-2009 caused by a decline in reservoir precipitously; many cities, including Sydney, Melbourne, and Perth, were affected. From 2001 to 2012, \$4.5 billion provided by the federal government for relief to drought-pretentious agriculturalists and small businesses. According to National Disaster Management Authority (NDMA) report November 2018 titled "Report on prevailing drought-like situation in Sindh with Particular reference to district Tharparkar" it is reported that during last four years District Tharparkar received very nominal and inadequate rains during monsoon season. A detailed statement of rainfall since 2014. This has adversely affected Livestock and Agriculture in 167 Deaths which are the main sources of livelihood in Tharparkar. Lower incomes resulted in health and nutrition issues in most vulnerable groups like pregnant/lactating women and children particularly newly born (NDMA, 2018). A recent study declares that more than 4500 drought events occurred during the period of 1951-2016 across the world, in which 52 events identified as megadrought events (Spinoni et al., 2019).

It is challenging to accurately ascertain the inception and end up of a drought event. Therefore, the appropriate knowledgeabout the commencement and ispersion of drought over a region is very effective for drought preparedness. However the assessment of drought is principal importance for the management of water resources

Table 1. Classification of drought based on SPI and RDI

Drought class	SPI & RDI
Extreme wet	2.0+
Severe wet	1.50 to 1.99
Moderate wet	1.00 to 1.49
Near normal	-0.99 to 0.99
Moderate drought	-1.00 to -1.49
Severe drought	-1.50 to -1.99
Extreme drought	≤-2.00

and planning (Mishra & Singh, 2010). This involves understanding historical drought in the area along their impact throughout their occurrence. Hence drought indices are key tool to gain useful information for drought monitoring in a region. Various drought indicators are being used around the world since climatic circumstances vary from terrain to terrain so just single drought index cannot deliver the comprehensive evidence about drought (Table 1). Most common and broadlyemployed index is standardized precipitation index (Spinoni et al.), developed by (McKee, Doesken, & Kleist, 1993) to examine and see drought situations as advised by (Adnan et al., 2018; Ahmad, Parvaze, Majid, & Kanth, 2016). Another index Reconnaissance Drought Index (RDI)is used by many European countriesand is also recommended by (Adnan et al., 2018) to supervise drought significance in Pakistan. Thus, the keyaim of the study is to assess thehistorical drought event in Tharparkar, Pakistan. This would be beneficial for national meteorological services, mitigation planners and water resources management to cope up with future climatic Change adaptation in Tharparkar.

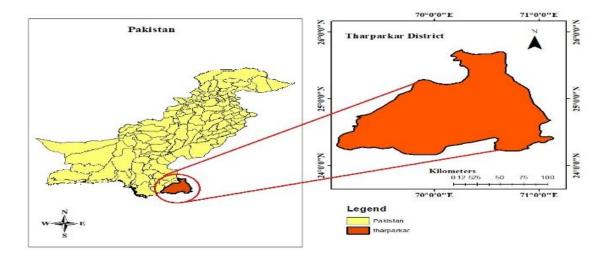


Fig. 1 Map of Pakistan showing Tharparkar District

Data and Methodology

Study area

Tharparkar district lies in the domain of 24°10′N-25°45′ N and 69° 04E-71°06E with an area of 19,638 km2 as shown in Figure.1.

Administratively, Tharparkar district is divided into seven talukas Chachro, Dhali, Mithi, Diplo, Kaloi, Islamkot, and Nagarparkar. It is composed of virtually two thousand small villages. The population of Tharparkar district is 1,649,661 estimated by 2017 census as shown in Figure.2. Average population growth rate of district Tharparkar is 3.15 estimated from 1998-2017 census.

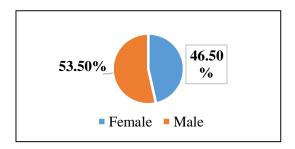


Fig. 2: Population Figure of Tharparkar district Census 2017 (statistics)

Data

The long term (1979-2013) daily climate data of two meteorological elements such as precipitation (mm) and temperature have been used in current study. For evaluating two indicators, 35 year's data were downloaded from Global Weather Data for SWAT.

Drought indices

Standardized Precipitation Index (Spinoni et al.)

Due to single input variable the SPI index has more importance around the world for drought identification. SPI use precipitation values for exclusive month-to-month timescales (3, 6, 9, 12, 24, and 48 months) to capture the

temporal performance of drought in the region. SPI required 20-30 years month-to-month records, 50-60 12 months (or extra) facts being most useful favored by (Guttman, 1994). The gamma distribution is found by (Thom, 1958 #20), to fit the climatological precipitation sequence. The gamma distribution is described with the assistance of its PDF, (probability density function).

$$g(x) = \frac{1}{\beta^a \Gamma(a)} x^{a-1} e^{-x/\beta}, \text{ for } x > 0 \cdots (1)$$

In which β and a are the scale and shape parameters respectively, $\Gamma(x)$ is the gamma function and, x is precipitation amount. Parameter β and a of the pdf are estimated for the different time scale. Where β and a estimation are given as:

$$\beta = \frac{\bar{x}}{a}, \ a = \frac{1}{4A} \left(1 + \sqrt{1 + \frac{4A}{3}} \right) \cdots (2)$$

Where
$$A = \ln(x) - \frac{\sum \ln(x)}{n}$$

n is the number of observations.

Reconnaissance Drought Index (RDI)

Reconnaissance Drought Index (RDI) is a meteorological index developed by (Tsakiris & Vangelis, 2005). The RDI is based totally each on potential evapotranspiration (PET) and cumulative precipitation (P), that is one calculated (PET) determinant and one measured (P). The initial value (λk) is based on the proportion between potential evapotranspiration (PET) and precipitation. The PET calculation method does not affect the RDI results (Vangelis, Tigkas, & Tsakiris, 2013). So, RDI is calculated by applying the following equations

$$\lambda_k^{(i)} = \frac{\sum_{j=1}^k P_{ij}}{\sum_{j=1}^k PET_{ij}}, i=1 \ N \text{ and } j=1 \ k \ \cdots (3)$$

Where PET_{ij} and P_{ij} are the potential evapotranspiration and precipitation of the j-th month of the i-th year and N is the total number of years of the available climate data. The standard RDI_{st} states that the λ_k takes the values of lognormal distribution given as:

$$RDI_{st}^{(i)} = \frac{y_k - \overline{y}_k}{\lambda_{yk}} \quad \cdots (4)$$

 y_k is the, $\overline{y_k}$ is the arithmetic mean and λ_{yk} is the standard deviation.

Result and discussion

The present study analysis the historical drought events for the period (1979-2013) in Tharparkar district. The annual time series is helpful to identify the drought event in the region. Therefore, the time series of SPI and RDI has been conduted for tharparkar as shown in Figure.3 and 4 respectively.

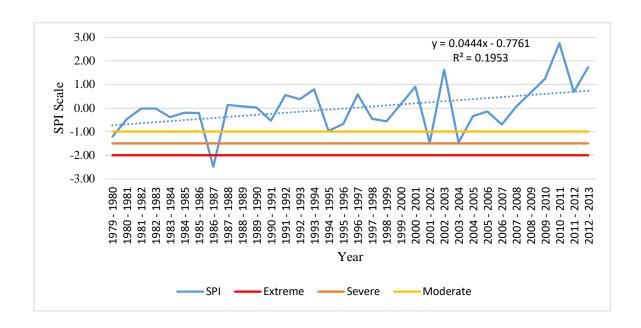


Fig. 3: Annual time series analysis of SPI in Tharparkar

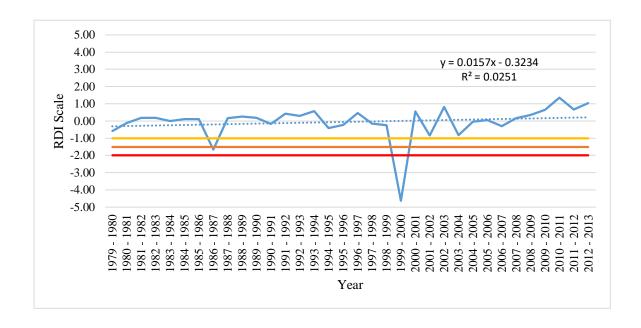


Fig. 4: Annual time series analysis of RDI in Tharparkar

According to the result obtained by SPI for tharparkar, the extreme drought event occurred in 1986-1987 where SPI value is -2.48. The severe drought captured by SPI is during the period 2001-2004 with -1.50 SPI value.

RDI figured out the extreme drought in the year 1999-2000, where RDI values reaches at -4.63. however the severe drought perceived in the year 1986-1987 with RDI value (-1.50). The moderate drought observed during the period 2001-2004 respectively. The main objective of the present study was to assess the historical drought by using the drought criterion SPI and RDI. Both the indices shows good competence to captured the drought in the selected region. The drought period 2001-2004 comprehend by both drought indices is also assessed by (Malik, Taylor, Szeto, & Khan, 2013). The correlation observed by SPI is (r=0.441), where the stronge correlation comprehend by RDI (r=0.158). The result obtained from the indicators pointed out that the indices SPI and RDI are better predictor to monitor the drought in the region. This study will be helpful to develop a mitigation plane and to tackle drought events and to build future scemarios of droght related desasters in the future.

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